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Atmos. Chem. Phys., 8, 3899-3917, 2008
www.atmos-chem-phys.net/8/3899/2008/
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Observations of  $HNO_3$ ,  $\Sigma AN$ ,  $\Sigma PN$  and  $NO_2$  fluxes: evidence for rapid  $HO_x$  chemistry within a pine forest canopy

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Abstract. Measurements of exchange of reactive nitrogen oxides between the atmosphere and a ponderosa pine forest in the Sierra Nevada Mountains are reported. During winter, we observe upward fluxes of NO<sub>2</sub>, and downward fluxes of total peroxy and peroxy acyl nitrates (SPNs), total gas and particle phase alkyl and multifunctional alkyl nitrates  $(\Sigma ANs_{(q+p)})$ , and the sum of gaseous HNO3 and semi-volatile NO3 particles (HNO3 (q+p)). We use calculations of the vertical profile and flux of NO, partially constrained by observations, to show that net midday  $\Sigma NO_{vi}$  fluxes in winter are -4.9 ppt m s<sup>-1</sup>. The signs and magnitudes of these wintertime individual and  $\Sigma \text{NO}_{vi}$  fluxes are in the range of prior measurements. In contrast, during summer, we observe downward fluxes only of  $\Sigma ANs_{(q+p)}$ , and upward fluxes of  $HNO_{3(q+p)'}$   $\Sigma PNs$  and  $NO_2$  with signs and magnitudes that are unlike most, if not all, previous observations and analyses of fluxes of individual nitrogen oxides. The results imply that the mechanisms contributing to NO<sub>v</sub> fluxes, at least at this site, are much more complex than previously recognized. We show that the observations of upward fluxes of  $HNO_{3(q+p)}$  and  $\sigma PNs$  during summer are consistent with oxidation of NO<sub>2</sub> and acetaldehyde by an OH x residence time of  $1.1 \times 10^{10}$  molec OH  $cm^{-3}$  s, corresponding to 3 to  $16 \times 10^7$  molecules  $cm^{-3}$  OH within the forest canopy for a 420 to 70 s canopy residence time. We show that  $\Sigma AN_{(q+p)}$ fluxes are consistent with this range in OH if the reaction of OH with  $\Sigma$ ANs produces either  $HNO_3$  or  $NO_2$  with a 6–30% yield. Calculations of NO fluxes constrained by the NO2 observations and the inferred OH indicate that NOx fluxes are downward into the canopy because of the substantial conversion of  $\text{NO}_{x}$  to  $\text{HNO}_{3}$  and  $\sigma\text{PNs}$  in the canopy. Even so, we derive that NO<sub>x</sub> emission fluxes of ~15 ng(N) m<sup>-2</sup> s<sup>-1</sup> at midday during summer are required to balance the  $\mathrm{NO}_{\rm X}$  and  $\mathrm{NO}_{\rm V}$  flux budgets. These fluxes are partly explained by estimates of soil emissions (estimated to be between 3 and 6 ng(N) m<sup>-2</sup> s<sup>-1</sup>). One possibility for the remainder of the NO<sub>x</sub> source is large HONO emissions. Alternatively, the 15 ng(N)  $m^{-2} s^{-1}$  emission estimate may be too large, and the budget balanced if the deposition of  $HNO_3$  and  $\sigma PNs$  is slower than we estimate, if there are large errors in either our understanding of peroxy radical chemistry, or our assumptions that the budget is required to balance because the fluxes do not obey similarity theory.

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Citation: Farmer, D. K. and Cohen, R. C.: Observations of HNO<sub>3</sub>, ΣAN, ΣPN and NO<sub>2</sub> fluxes: evidence for rapid HO<sub>x</sub> chemistry within a pine forest canopy, Atmos. Chem. Phys., 8, 3899-3917, 2008. ■ <u>Bibtex</u> ■ <u>EndNote</u> ■ <u>Reference Manager</u>