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- Title and Author Search

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Correlation of pre-earthquake electromagnetic signals with laboratory and field rock experiments

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Abstract. Analysis of the 2007 *M*5.4 Alum Rock earthquake near San Jose, California showed that magnetic pulsations were present in large numbers and with significant amplitudes during the 2 week period leading up to the event. These pulsations were 1–30 s in duration, had unusual polarities (many with only positive or only negative polarities versus both positive and negative) and were different than other pulsations observed over 2 years of observation prior to the earthquake. The pulse sequence was sustained over a 2 week period prior to the earthquake, and then disappeared shortly after the quake. A search for underlying physics process that might explain these pulses was widely undertaken, and one theory (Freund, 2002) demonstrated that charge carriers were released when various types of rocks were stressed in a laboratory environment. It was also significant that the observed current carrier generation was transient, and resulted in pulsating current patterns. In an attempt to determine if this phenomenon occurred in the laboratory environment, the authors scaled up the physics experiment from a relatively small rock sample in a dry laboratory setting to a large 7 metric tonne boulder comprised of Yosemite granite. This boulder was located in a natural, humid (above ground) setting at Lake Tahoe, Ca. The boulder was instrumented with two Zonge Engineer Model ANT4 induction type magnetometers, two Trifield Air Ion Collector surface charge detector, a geophone, a Bruker Model EM27 Fourier Transform Infra Red (FTIR) spectrometer with Sterling cycle cooler, various temperature sensors. The boulder was stressed over about 5 h using expanding concrete (Bustartm), until it fractured into three main pieces. The recorded data showed surface charge build up, magnetic pulsations, impulsive air conductivity changes, and acoustical cues about 5 h before the boulder actually broke. These magnetic and conductivity pulse signatures resembled both the laboratory rock stress test results and the 30 October 2007 *M*5.4 Alum Rock earthquake field

The second part of this paper examined other California earthquakes prior to the Alum Rock earthquake, to see if magnetic pulsations were present prior to those events. A search for field examples of medium magnitude earthquakes was performed to identify earthquakes where functioning magnetometers were present within 20 km, the expected detection range of the magnetometers. Two earthquakes identified in the search were the 12 August 1998 *M*5.1 San Juan Bautista (Hollister Ca.) earthquake

the 28 September 2004 $M6.0$ Parkfield Ca. earthquake. Both of the sets were recorded using EMI Corp. Model BF4 induction magnetometer installed in equipment owned and operated by UC Berkeley. Unfortunately, no air conductivity or IR data were available for these earthquake examples. This new analysis of old data used the raw time series (10 samples per s), and examined the data for short duration pulsations that exceeded the normal background noise levels at each site, similar to the technique used at Alum Rock. Analysis of Hollister magnetometer, positioned 2 km from the epicenter, showed a significant increase in magnetic pulsations above quiescent threshold levels several weeks and especially 2 days prior to the quake. The pattern of positive and negative pulsations observed at Hollister, were similar, but not identical to Alum Rock in that the pattern of pulsations were interspersed with pulsation trains, and did not start 2 weeks prior to the quake, but 2 days prior. The Parkfield data (magnetometer positioned 19 km from epicenter) showed much smaller pre-earthquake pulsations, but the site had significantly higher conductivity (which attenuates the signals). An interesting fact was that significant pulsations occurred between aftershock sequences of quakes as the crustal stress patterns were migrating.

Comparing laboratory, field experiments with a boulder, and earthquake events, striking similarities were noted in magnetic pulsations and conductivity changes, as well as IR signals (where instrumented). Additional earthquake samples, taken with the appropriate detectors and within 15 km proximity to large ($>M5$) earthquakes, are still needed to provide more evidence to understand the variability between earthquakes and various electromagnetic signals detected prior to large earthquakes.

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