Estimating the Aggregate/Intraaggregate Mass Ratio of a Shrinking Soil

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Abstract: A recently introduced parameter, the ratio of an aggregate solid mass to a solid mass of an intraaggregate matrix (K ratio) is connected with the mean thickness of a deformable, but non-shrinking surface layer of aggregates and is a fundamental property of aggregated soils that essentially influences their shrinkage. The objective of this work is to suggest and validate an approach to estimating the K ratio at any soil clay content through characteristics of soil texture and structure. We derive an equation that reflects the interrelation between the K ratio and soil texture and structure. The K ratio can be estimated as the solution of the equation and is determined by the mean size of soil solids and the maximum size of soil aggregates in the oven-dried state, independently of a measured shrinkage curve. To validate the approach we use available data for eight soils

INTRODUCTION

A recently proposed model [1, 2] showed how a clay shrinkage curve under the influence of a silt-sand admixture as well as inter- and intraaggregate structures, is transformed into the soil shrinkage curve at any soil clay content. In general, the shrinkage curve of a soil is non-single valued since the crack volume contribution to the shrinking soil volume depends on sampling, sample preparation, sample size, and drying regime. The above model, unlike previous ones (e.g., [3-6]) investigated the reference shrinkage curve that, by definition, corresponds to shrinkage without interaggregate cracking and can be predicted in a single valued manner. In addition, the previous models justify some a priori taken mathematical expression for a shrinkage curve (different for each different approach) by the fitting of its parameters (from 3 to 11 depending on the approach) to experimental shrinkage curve data. Unlike these models, Chertkov [1, 2] derived the reference shrinkage curve of a soil from the shrinkage curve of a clay [7, 8] contributing to the soil and two new features of the intraaggregate soil structure (Fig. 1): (i) the existence and dewatering of a deformable, but nonshrinking aggregate surface layer (interface layer); and (ii) the existence and volume increase of intraaggregate lacunar pores at soil shrinkage. The interface layer exists at any soil clay content (Fig. 1). Its mean thickness increases with a clay content decrease. The lacunar pores exist at a clay content, c lower than a critical value, c^* (Fig. 1).

The derivation of the reference shrinkage curve explains the transition from a clay shrinkage curve to a soil shrinkage curve and leads to the understanding of the origin of the shape of a soil shrinkage curve. In addition, the reference shrinkage curve can be used in estimating the contribution of the crack volume to the soil shrinkage curve, the soil hydraulic properties, and in other applications.

The methodology of the reference shrinkage curve prediction was based on detailed accounting for possible contributions to the soil volume and water content during shrinkage. The reference shrinkage curve is determined by eight physical soil parameters: oven-dried specific volume (Y_z) , maximum swelling water content (W_h) , mean solid density (ρ_s) , soil clay content (c), oven-dried structural porosity (P_z) , the ratio of aggregate solid mass to solid mass of intraaggregate matrix (K), the lacunar factor that characterizes the rate of the lacunar pore volume change with water content (k), and oven-dried lacunar pore volume (U_{lpz}) . The model [1, 2] was validated using data for six soils from [9], a soil from [10], and eight soils from [11].

The last three parameters (K, k, and $U_{\rm lpz}$) present new fundamental physical soil properties and as such, in principle, can be measured or estimated from some measurements independently of a measured shrinkage curve. The methods of estimating k and $U_{\rm lpz}$ were briefly indicated [2]. Estimating the K ratio independently of a measured shrinkage curve in the particular case of the thin interface layer (Fig. 1 at $c>c^*$ when the lacunar pores are lacking) was noted without validation [12]. The objective of this work is to propose and validate an approach for estimating the K ratio in the general case of any soil clay content, from the oven-dried aggregate-size distribution and soil texture. Notation is summarized at the end of the paper.

THEORY

General Expression for the K Ratio

The *K* ratio can be written as [1, 2]

$$K = (1 - U_i/U_h)^{-1} \tag{1}$$

where U_i is the specific volume of the deformable, but nonshrinking aggregate surface layer (interface layer) (Fig. 1), and U_h is the maximum specific volume of the intraaggregate matrix (Fig. 1) when shrinkage starts (at the total water content $W=W_h$). Each aggregate has the interface layer (Fig. 1) that contributes to U_i . This means that U_i can be expressed through the aggregate-size distribution at the total water content $W=W_h$ when shrinkage starts as

$$U_{i} = \frac{(1 - P_{h})}{\rho_{h}} \int_{X_{\min}}^{X_{m}} \frac{V_{i}(X, x)}{V_{a}(X)} \frac{dF(X, P_{h})}{dX} dX .$$
 (2)

In Eq.(2) $F(X,P_h)$ is the cumulative aggregate-size distribution as a function of the current aggregate size, X and

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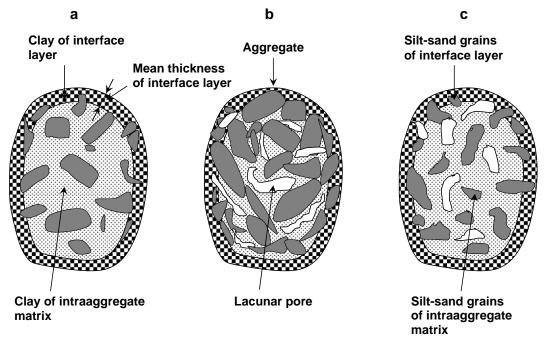


Fig. (1). The illustrative scheme of the internal structure of aggregates at a clay content (reproduced from [2], Fig. 2): (a) $c > c^*$, without lacunar pores; (b) $c < c^*$, with lacunar pores and silt-sand grain contacts at W=0; and (c) $c < c^*$, with lacunar pores and non-totally contacting silt-sand grains at any water content. c^* is the critical soil clay content.

structural porosity, P_h of the soil at $W=W_h$; $V_i(X, x)$ is the volume of the interface layer of the mean thickness, x relating to a separate aggregate of size X (Fig. 1); $V_a(X)$ is the total volume of a separate aggregate of size X; X_{min} and X_m are the minimum and maximum sizes of aggregates, respectively, at $W=W_h$; ρ_b is the dry bulk density of the soil (the soil solid mass per unit soil volume at $W=W_h$).

 $(dF(X,P_h)/dX) dX$ in Eq.(2) gives a volume of aggregates in the size range dX per unit volume of aggregates of all the possible sizes, $X_{\min} \le X \le X_{\min}$ (at $W = W_h$). $(1/V_a(X))(dF(X,P_h)/V_h)$ dX) dX in Eq.(2) gives the number of aggregates in the size range dX per unit volume of all aggregates (at $W=W_h$). ($V_i(X, Y_h)$). $(x)/V_a(X)(dF(X,P_h)/dX) dX$ in Eq.(2) gives the interface layer volume (of the mean thickness x) relating to aggregates in the size range dX per unit volume of all aggregates (at $W=W_h$). Integration over X gives the total interface layer volume (of the mean thickness x) of all aggregates per unit volume of the aggregates (at $W=W_h$). Multiplication by $(1-P_h)$ gives the total interface layer volume of all aggregates per unit volume of the soil including structural pores (at $W=W_h$). Finally, division by ρ_b gives the total interface layer volume per unit mass of oven-dried soil, that is, the specific volume U_i of the interface layer.

Replacement of U_i in Eq.(1) with U_i from Eq.(2) leads to the general expression for the K ratio. In the following sections we specify the different values entering the K ratio expression from Eq.(2).

Transformation of the Bulk Density Term

One can see that $U_{\rm h}/(1-P_{\rm h})$ is the maximum specific soil volume at $W=W_{\rm h}$ when the interface layer and intraaggregate matrix (Fig. 1) have the same water content and specific volume [1, 2]. This means that the multiplier $(1-P_{\rm h})/(\rho_{\rm b}U_{\rm h})$ before the integral in the above expression for the K ratio is reduced to unity

$$(1-P_h)/(\rho_b U_h)=1$$
 . (3)

Approximation of Aggregate Shape

The aggregate-size distribution, $F(X,P_h)$ (see the following section), as any size distribution (of pores, grains, aggregates, and others), by definition, only accounts for the size of aggregates (X), but not their shape. Therefore, considering the $V_i(X,x)/V_a(X)$ ratio from Eq.(2) we can in the same approximation neglect the effect of the aggregate shape on its volume (V_a) and volume (V_i) of its interface layer of the thickness X (Fig. 1). This means that when specifying the V_i/V_a ratio in Eq.(2), we can consider in the same approximation the aggregates of the simplest shape, that is, the spheres of diameter X. One can easily be convinced that for a sphere of diameter X and its surface layer of thickness X the V_i/V_a ratio should be

$$V_i(X, x)/V_a(X)=1-(1-2x/X)^3$$
 (4)

In addition, the same V_i/V_a ratio holds for a cube of side X and its surface layer of thickness x. In general, the (mean) interface layer thickness, x can have any value in the range 0 < x < X/2 depending on the soil clay content. In the case of the clay content higher than critical $(c > c^*)$ we deal with the relatively thin interface layer (x < < X/2) when the K ratio is in the range $1 < K < \equiv 1.3$ [1]. In this case Eq.(4) gives $V_i/V_a \cong 6x/X$. Estimation of the K ratio at $c > c^*$ (without experimental validation) was briefly considered [12]. At a clay content lower than critical $(c < c^*)$ the non-shrinking interface layer can even nearly occupy the total volume of an aggregate $(x \rightarrow X/2)$ when the clay content is sufficiently small, in this case $K \rightarrow \infty$

Aggregate-Size Distribution

Based on the intersecting surfaces approach to soil structure [13] we use in Eq.(2) the simplest aggregate-size distribution, $F(X, P_h)$ as

$$F(X, P_h) = F(\eta, P_h) = (1 - P_h^{Io(\eta)/8.4})/(1 - P_h)$$
 (5)

where

$$\eta = (X - X_{\min}) / (X_{\mathrm{m}} - X_{\min}), \qquad X_{\min} \le X \le X_{\mathrm{m}} \tag{6}$$

and

$$I_0(\eta) = \ln(6)(4\eta)^4 \exp(-4\eta)$$
 (7)

In the particular, but important, case of an aggregated soil with negligible structural porosity $(P_h \rightarrow 0)$ the $F(X,P_h)$ distribution in Eq.(5) should be replaced with [13]

$$F(\eta, 0) = \{1 - \exp[-I_0(\eta)]\} / [1 - \exp(-8.4)]. \tag{8}$$

Structural Porosity

The initial structural porosity, P_h (i.e., at $W=W_h$) that enters the $F(X,P_h)$ distribution (Eq.(5)) can be calculated as [12]

 $P_{\rm h} = [1 + \{ [(u_{\rm h}/u_{\rm s})/(\rho_{\rm s}K) + U_{\rm i}] \}$

$$/[(u_z/u_s)/(\rho_s K) + U_i] \{ (1 - P_z)/P_z \}^{-1}, \quad 0 \le W \le W_h$$
 (9)

where u_s is the relative volume of solids in the intraggregate matrix; u_z is the relative oven-dried volume of the intraaggregate matrix; and u_h is the relative volume of the intraaggregate matrix at $W=W_h$. The values of u_s , u_z , and u_h as well as U_i and, therefore, P_h (Eq.(9)) are calculated through the above input parameters: Y_z , W_h , ρ_s , c, P_z , K, k, and U_{lpz} [1, 2]. Thus, P_h is a function of the K ratio (at given parameters Y_z , W_h , ρ_s , c, P_z , k, and U_{lpz}).

The Relation between the Mean Thickness of an Interface Layer (x) and Soil Texture

The soil of the interface layer and that of an intraaggregate matrix (Fig. 1) are characterized by the same texture, i.e., by the same weight fractions of clay particles, silt, and sand grains. We believe that the mean thickness of the interface layer, at least by order of magnitude, is directly connected with a certain characteristic size of internal nonhomogeneities of the intraaggregate soil matrix (Fig. 1). The natural (and single) pretenders for the role of such nonhomogeneities (at the scale we consider, Fig. 1) are different solids (clay particles, silt, and sand grains), and the single natural candidate for the role of such characteristic size is the mean size of the soil solids (non-homogeneities) by their weight fractions, x_n . Thus, we accept the natural assumption - that the doubled mean interface layer thickness (2x) is approximately equal to the mean size of the soil solids (x_n) :

$$2x \cong x_n$$
. (10)

At high clay content $(c\rightarrow 1)$ x_n strives to a clay particle size, and the interface layer (of the x thickness; Fig. 1) is very thin. That is, $K \rightarrow 1$. At small clay content ($c << c^*$) the mean size of soil solids, x_n strives to a silt-sand grain size (depending on the grain-size distribution). With that aggregates are also degenerated to rigid silt and sand grains. This means that the intraaggregate matrix (Fig. 1) disappears, and the aggregate /intraaggregate mass ratio, $K \rightarrow \infty$.

Minimum and Maximum Sizes of Aggregates

 X_{\min} and X_{\min} in Eq. (2) and (6) relate to $W=W_{h}$ and can be expressed through $X_{\min z}$ and $X_{\max z}$ values in oven-dried state, respectively. At a given water content the intraaggregate matrix volume is proportional both to the cube of the difference between the current aggregate size and doubled interface layer thickness, $(X(W)-2x)^3$ (Fig. 1) and to the relative volume of the intraaggregate matrix, u(W). In particular, for the maximum aggregate size, $X_{\rm m}$ at $W=W_{\rm h}$ and that, $X_{\rm mz}$ at *W*=0 one can write

$$(X_{\rm m}-2x)^3 \propto u(W_{\rm h}) \equiv u_{\rm h} \tag{11}$$

and

$$(X_{\text{mz}}-2x)^3 \propto u(0) = u_z$$
, (12)

respectively. It follows that

$$X_{\rm m} = 2x + (X_{\rm mz} - 2x)(u_{\rm h}/u_{\rm z})^{1/3} . {13}$$

The above consideration is applicable to any aggregate size. Replacing in Eq.(11)-(13) the maximum aggregate size values, $X_{\rm m}$ and $X_{\rm mz}$ with the minimum aggregate size values, X_{\min} and X_{\min} , respectively, we obtain

$$X_{\min} = 2x + (X_{\min}z^{-2}x)(u_h/u_z)^{1/3}$$
 (14)

The maximum aggregate size in the oven-dried state $(X_{\rm mz})$ can essentially exceed the doubled mean interface layer thickness (2x) (Fig. 1). Unlike that the difference between X_{minz} and 2x in Eq.(14) in any case should be relatively small. That is $(X_{\text{minz}}-2x)/(2x) << 1$. Indeed, any current aggregate size (X) of a shrinking soil should exceed the doubled mean interface layer thickness (2x) (Fig. 1; X>2x). Otherwise, aggregates would turn into rigid formations. That is, they would be degenerated to silt and sand grains. It follows that the minimum aggregate size [at any water content; e.g., both X_{\min} and $X_{\min z}$, see Eq. (14)] should be close to 2x (but exceeding that). Thus, Eq. (14) can be rewritten as

$$X_{\min} \cong X_{\min} \cong 2x . \tag{15}$$

Equations (10) and (15) lead to a simple, but important structural relation between the minimum aggregate size (X_{\min}) and the mean size of the soil solids - nonhomogeneities (x_n) as

$$X_{\min} \cong X_{\min} \cong X_{n}$$
 (16)

In addition, Eq.(10) and (13) present the maximum aggregate size $(X_{\rm m})$ as

$$X_{\rm m} = x_{\rm n} + (X_{\rm mz} - x_{\rm n})(u_{\rm h}/u_{\rm z})^{1/3}$$
(17)

It is worth noting that Eq.(16)-(17) relate to any soil, and permit one to trace the transition from a shrinking to a rigid soil. Indeed, if the shrinkage of a soil is negligible (the clay content, $c \rightarrow 0$) the ratio, $u_h/u_z \rightarrow 1$ and Eq.(17) gives $X_m \rightarrow X_{mz}$ as it should be for a rigid soil. In addition, in this case $(c\rightarrow 0)$ x_n is reduced to the mean size of silt and sand grains. Therefore, the minimum aggregate size, X_{\min} (Eq.(16)) is reduced to the mean size of silt and sand grains (x_n) , and the maximum aggregate size (X_m) is obviously reduced to the maximum sand grain size.

Equation for the K Ratio and General Algorithm of its Numerical Solution

Using the above presentations of $(1-P_h)/(\rho_b U_h)$ (Eq.(3)), V_i/V_a (Eq.(4)), $F(X,P_h)$ (Eq.(5)-(8)), and X_{min} (Eq.(16)) one can present Eq.(1) and (2) as

$$K = \left\{ (1 - \alpha)^3 - 3[\alpha/(1 - \alpha)] \int_{0}^{1} \frac{\eta^2 F(\eta, P_h) d\eta}{0 (\eta + [\alpha/(1 - \alpha)])^4} \right\}^{-1}$$
(18)

where

$$\alpha = x_n/X_m$$
 (19)

According to the physical meaning of x_n and X_m (see above) the α ratio can take values in the range 0<a<1. The case of $\alpha \rightarrow 0$ (in practice, $0 < \alpha <<1$) corresponds to high clay contents $(c \rightarrow 1)$. Indeed [as noted above, see text after Eq. (10)] at high clay content x_n strives to a clay particle size. That is, $x_n/X_m <<1$ and, therefore, $\alpha <<1$ (Eq. (19)). One can see that in this case (when $\alpha \rightarrow 0$), according to Eq.(18), $K \rightarrow 1$ as it should be at $c \rightarrow 1$ [2]. The opposite case of $\alpha \rightarrow 1$ (but less than 1) corresponds to low clay contents ($c\rightarrow 0$, in practice $c << c^*$). Indeed [as noted above, see text after Eq.(10)] at low clay content x_n strives to a silt-sand grain size, and X_m strives to the maximum sand grain size. That is, the maximum possible value of the x_n/X_m ratio is unity (at a special possible grain-size distribution), and, therefore, $\alpha \rightarrow 1$ (Eq. (19)). One can see that in this case (when $\alpha \rightarrow 1$), according to Eq. (18), $K \rightarrow \infty$ (in practice K >> 1) as it should be at $c \rightarrow 0$ and transition to a rigid soil [2].

Accounting for Eq. (16) for $X_{\rm m}$ one can write α (Eq. (19)) as

$$\alpha = [1 + (X_{\text{mz}}/x_{\text{n}} - 1)(u_{\text{h}}/u_{\text{z}})^{1/3}]^{-1} . \tag{20}$$

The u_h/u_z ratio at given parameters Y_z , W_h , ρ_s , c, P_z , k, and U_{lpz} depends on K [2]. This means that $\alpha = \alpha(K)$ from Eq.(20) is a known function of K. In addition, $P_h = P_h(K)$ (Eq.(9)). Thus, Eq.(18) is one relative to K as an unknown value. Given the above seven input parameters: Y_z , W_h , ρ_s , c, P_z , k, and U_{lpz} and two additional input parameters, X_{mz} and x_n (instead of K) one can solve Eq.(18) to find the K ratio of the shrinking soil. Note, that additional parameters (X_{mz} and x_n) are generalized characteristics of aggregate sizes (X_{mz}) and soil texture (x_n).

In a practical calculation of K from Eq.(18) with given Y_z , W_h , ρ_s , c, P_z , k, U_{lpz} and additional parameters, X_{mz} and x_n , the simplest general algorithm is as follows. One can increase a test K value with a step ΔK (e.g., 0.001) starting from K=1. Using the test K value together with given Y_z , W_h , ρ_s , c, P_z , k, U_{lpz} one can find u_s , u_z , u_h , and U_i [2], then P_h (Eq.(9)), α (Eq.(20), using the additional parameters, X_{mz} and x_n), and, finally, the right part of Eq.(18) that we designate (at the test K value) as K'(K). Such calculation of the K'(K) values is continued with increasing K by step ΔK until K<K'(K). If K_1 < $K'(K_1)$ and K_2 = K_1 + ΔK > $K'(K_2)$ the true K value (i.e., the K root of Eq.(18)) is K_1 + ΔK /2. Decreasing ΔK

one can increase the accuracy of the K calculation. Thus, the K ratio of a shrinking soil can be calculated from other input parameters, Y_z , W_h , ρ_s , c, P_z , k, U_{lpz} and generalized characteristics of aggregate sizes, X_{mz} and soil texture, x_n .

MATERIALS AND METHODS

Two Possible Ways to Validate the K Ratio Estimation

To validate the approach to K ratio estimation as applied to a soil, one needs the data set, Y_z , W_h , ρ_s , c, P_z , K, k, and U_{lpz} (including K) that permits one to predict the soil reference shrinkage curve [1, 2], and, in addition, data on the soil texture and structure to independently estimate the mean size of soil solids, x_n and the maximum soil aggregate size in the oven-dried state, X_{mz} . Then, two ways of the validation are possible.

The first, direct way includes the following steps:

- (i) estimating the soil parameters u_s , u_z , u_h , and U_i that correspond to the given data set, Y_z , W_h , ρ_s , c, P_z , K, k, and U_{lpz} according to [1, 2];
- (ii) estimating P_h (Eq.(9)) and α (using Eq.(20) with the given x_n and X_{mz} values);
- (iii) numerical solving Eq.(18) relative to K with $F(\eta, P_h)$ from Eq.(5)-(8) and at found P_h and α (for the general algorithm see the end of the previous section); and
- (iv) comparison between the found K root and the initially given K ratio (that enters the data set to predict the reference shrinkage curve of the soil).

The second, opposite, way includes the following steps:

- (i) this step coincides with (i) of the first way;
- (ii) estimating P_h (Eq.(9));
- (iii) numerical solving Eq.(18) relative to α with $F(\eta,P_h)$ from Eq.(5)-(8) and at the given K ratio;
- (iv) estimating the $\alpha' = x_n/X_{mz}$ ratio from Eq.(20) with the u_b/u_z ratio (see step (i)) and found α as

$$\alpha' = x_{n}/X_{mz} = \{1 + [(1-\alpha)/\alpha](u_{h}/u_{z})^{-1/3}\}^{-1};$$
(21)

(v) comparison between the above found $\alpha' = x_n/X_{mz}$ and the ratio x_n/X_{mz} corresponding to the initially and independently given x_n and X_{mz} values.

Data Used

Currently, examples of the data sets for the reference shrinkage curve prediction (Y_z , W_h , ρ_s , c, P_z , K, k, and U_{lpz}) are only available from [1, 2]. These data sets relate to the experimental shrinkage curves of six soils from [9], a soil from [10], and eight soils from [11]. However, data that are connected with the x_n and X_{mz} values of the soils only are in [11]. For this reason we only consider the eight soils from [11]. Table 1 reproduces the input data (Y_z , W_h , ρ_s , c, P_z , K, k, and U_{lpz}) of the eight soils for the reference shrinkage curve prediction from [1, 2]. In addition, Table 1 contains the predicted parameters, u_s , u_z , u_h , and U_i for the soils from [1, 2].

Work [11] contains the data for the eight soils on weight fractions $(p_1, p_2, p_3, p_4, p_5)$ of soil solids from five size ranges. Table 2 reproduces the p_1, p_2, p_3, p_4, p_5 values. These

Table 1. Input Parameters of the Eight Soils for Reference Shrinkage Curve Prediction and Predicted Parameters Participating in the K ratio Estimation[†]

Data Source		Soil Number	Input Parameters from [11]				Input Parameters of New Soil Charac- teristics from [1, 2]			Predicted Parameters from [1, 2]				
			Yz	$W_{ m h}$	$\rho_{\rm s}$	c	Pz	K	k	$U_{ m lpz}$	$u_{\rm s}$	u_z	u_{h}	U_{i}
			dm³ kg-¹	kg kg ⁻¹	g cm ⁻³					dm³ kg-¹				dm³ kg ⁻¹
Ferruginous soil from Fig. (3) of [11]	A horizon	1	0.642	0.200	2.337	0.065	0	3	0.929	0.153	0.516	0.766	0.778	0.430
	B1 horizon	2	0.692	0.235	2.337	0.340	0.038	1.894	0.387	0.040	0.476	0.724	0.760	0.322
	B2 horizon	3	0.688	0.211	2.337	0.353	0.051	1.538	0.524	0.050	0.503	0.760	0.784	0.233
	AB horizon	4	0.682	0.229	2.337	0.187	0	2.499	0.771	0.130	0.483	0.751	0.782	0.416
Ferralitic soil from Fig. (3) of [11]	A horizon	5	0.878	0.543	2.608	0.523	0	1.892	0.523	0.132	0.261	0.562	0.638	0.442
	B1 horizon	6	0.696	0.351	2.608	0.648	0	1.433	0.340	0.028	0.354	0.627	0.677	0.222
	B2 horizon	7	0.731	0.362	2.608	0.648	0.031	1.424	0.121	0.0074	0.346	0.625	0.673	0.222
	AB horizon	8	0.714	0.372	2.608	0.606	0	1.329	0	0	0.340	0.624	0.670	0.187

[†] Y₂, specific volume of an oven-dried soil; W_h, total gravimetric water content at maximum aggregate and soil swelling; ρ₈, mean density of solids; c, weight fraction of clay solids; Pz, oven-dried structural porosity of the soil; K, experimental estimate of aggregate/intraaggregate mass ratio; k, experimental estimate of lacunar factor value; U_{1pz} , experimental estimate of oven-dried value of the specific volume of lacunar pores; u_s , relative volume of all solids; u_z , relative volume of intraaggregate matrix at its shrinkage limit; u_h , relative volume of intraaggregate matrix at maximum swelling point; U_h , contribution of the interface aggregate layer to the specific volume of soil aggregates.

data indicate the size range of the largest sand grains as >0.2 mm. The maximum sand grain size, $x_{\rm m}$ is unavailable for the soils. Thus, the data per se are not sufficient to find the accurate x_n and X_{mz} values and, correspondingly, the accurate x_n/X_{mz} ratio and, then, α from Eq.(20). For this reason we could not immediately use both of the above possible ways to validate the K ratio estimation. Nevertheless, the data from Table 2 contain the sufficient information to obtain approximate x_n and X_{mz} values. We could use these approximate estimates of x_n and X_{mz} in frames of the second way after its modification.

Modified Second Way to Validate the K Ratio Estimation

The modified second way differs from the initial one only with respect to step (v) (see the text after Eq.(21)). In the modified variant we compare the found $\alpha' = x_n/X_{mz}$ (from Y_z , W_h , ρ_s , c, P_z , K, k, and U_{lpz}) not with an independent estimate of the x_n/X_{mz} ratio that flows out of the soil texture and structure, but with upper and lower estimates of the ratio. Using the texture data from Table 2 and some reasonable considerations we can find these upper and lower estimates of the x_n/X_{mz} ratio for the eight soils (instead of an accurate estimate for each soil).

The upper estimate of the x_n/X_{mz} ratio for each soil was obtained as follows. We exclude the clay fraction, <0.002 mm (Table 2). That gives the upper estimate of the mean solids size, x_n as

$$x_{\text{n up}} = \{0.011p_2 + 0.35p_3 + 0.125p_4 + [0.2 + (x_{\text{m}} - 0.2)]p_5\}/(p_2 + p_3 + p_4 + p_5)$$
(22)

(0.011, 0.35, and 0.125 mm are the mean sizes of 0.002-0.02, 0.02-0.05, and 0.05-0.2 mm ranges). Accounting for the maximum sand grain size, $x_{\rm m}$ being somewhere between 1 and 2 mm (e.g., [14]), we take three possible values, 1, 1.5, and 2 mm for $x_{\rm m}$ in Eq.(22). Since $x_{\rm n up}$ from Eq.(22) corresponds to a silt and sand mixture (without clay particles) the maximum aggregate size $X_{\rm mz}=x_{\rm m}$. Then, the $x_{\rm n~up}/x_{\rm m}$ ratio gives the upper estimate of the real x_n/X_{mz} ratio (that corresponds to the initial soil with clay) because the clay addition decreases the x_n value compared to $x_{n \text{ up}}$ and increases the X_{mz} value compared to $x_{\rm m}$. That is, the upper estimate, $(x_{\rm n}/X_{\rm mz})_{\rm up}$, of the x_n/X_{mz} ratio is

$$(x_n/X_{mz})_{up} = x_{n up}/x_m \tag{23}$$

with $x_{\text{n up}}$ from Eq. (22) and possible $x_{\text{m}}=1$, 1.5, and 2 mm.

Table 2. Data on the Texture of the Eight Soils from [11]

Soil Number	p_1	p_2	<i>p</i> ₃	p_4	p ₅
1	0.065	0.063	0.245	0.490	0.138
2	0.340	0.097	0.203	0.288	0.089
3	0.353	0.112	0.189	0.239	0.120
4	0.187	0.088	0.228	0.392	0.112
5	0.523	0.163	0.089	0.142	0.083
6	0.648	0.080	0.108	0.116	0.047
7	0.648	0.080	0.108	0.116	0.047
8	0.606	0.128	0.071	0.146	0.049

 $^{^\}dagger$ p_1 , p_2 , p_3 , p_4 , p_5 , weight fractions of soil solids from size ranges, <0.002, 0.002-0.02, 0.02-0.05, 0.05-0.2, and >0.2 mm, respectively

The lower estimate of the x_n/X_{mz} ratio for each soil was obtained as follows. We exclude the sand fraction, >0.2 mm (Table 2). That gives the lower estimate of the mean solids size, x_n as

$$x_{\text{n low}} = (0.001p_1 + 0.11p_2 + 0.35p_3 + 0.125p_4)$$

$$/(p_1 + p_2 + p_3 + p_4) . \tag{24}$$

As an upper estimate of $X_{\rm mz}$ we can take the same value as above, $X_{\rm mz}=x_{\rm m}$ (=1, 1.5, and 2 mm). Then, the $x_{\rm n~low}/x_{\rm m}$ ratio gives the lower estimate of the real $x_{\rm n}/X_{\rm mz}$ ratio (that corresponds to the initial soil with the sand fraction >0.2 mm) since the addition of the sand fraction increases the $x_{\rm n}$ value compared to $x_{\rm n~low}$ and can decrease $X_{\rm mz}$ compared to the $x_{\rm m}$ value. That is, the lower estimate, $(x_{\rm n}/X_{\rm mz})_{\rm low}$, of the $x_{\rm n}/X_{\rm mz}$ ratio is

$$(x_n/X_{mz})_{low} = x_{n low}/x_m$$
 (25)

with $x_{n \text{ low}}$ from Eq.(24) and possible $x_m=1, 1.5, \text{ and } 2 \text{ mm}$.

If the found $\alpha' \equiv x_n/X_{mz}$ (Eq. (21)) gets into the interval between the independent lower and upper estimates $((x_n/X_{mz})_{low})$ and $(x_n/X_{mz})_{up}$) this means that the approach to the K ratio estimation is in compliance with available data and, in any case, does not contradict them.

RESULTS AND DISCUSSION

Table 3 shows the $\alpha = x_n/X_m$ and $\alpha' = x_n/X_{mz}$ values found for the eight soils according to the second way (points (iii) and (iv), respectively). Table 3 also shows the lower and upper estimates of the x_n/X_{mz} ratio (Eq.(22)-(25)) that were found for the eight soils with x_m =1, 1.5, and 2 mm and using the p_1 , p_2 , p_3 , p_4 , p_5 fractions from Table 2. In addition, Table 3 shows P_h values relating to $W=W_h$, to illustrate the differ-

Table 3. Predicted α and α' Ratios, Independent Lower and Upper Estimates of the α' Ratio Found from Soil Texture[†], and the Structural Porosity, P_h at the Maximum Swelling Point

Soil Number	α≡x _n /X _m	$\alpha' = x_n/X_{mz}$	x _m mm	$(x_{\rm n}/X_{\rm mz})_{\rm low}$	$(x_{\rm n}/X_{\rm mz})_{\rm up}$	$P_{ m h}$
			1	0.0818	0.1638	
1	0.0994	0.0999	1.5	0.0545	0.1338	0
			2	0.0409	0.1188	
			1	0.0480	0.1441	
2	0.0725	0.0736	1.5	0.0320	0.1180	0.037
			2	0.0240	0.1049	
			1	0.0426	0.1662	
3	0.0485	0.0490	1.5	0.0284	0.1411	0.050
			2	0.0213	0.1286	
		0.0818	1	0.0650	0.1526	
4	0.0808		1.5	0.0433	0.1245	0
			2	0.0325	0.1105	
			1	0.0253	0.1519	
5	0.0538	0.0560	1.5	0.0169	0.1303	0
			2	0.0126	0.1195	
			1	0.0208	0.1349	
6	0.0288	0.0296	1.5	0.0139	0.1123	0
			2	0.0104	0.1009	
			1	0.0208	0.1349	
7	0.0375	0.0384	1.5	0.0139	0.1123	0.029
			2	0.0104	0.1009	
8	0.0224	0.0220	1.5	0.0159	0.1079	0
8	0.0224	0.0230	2	0.0120	0.0965	0

 $^{^{\}dagger}\alpha$, predicted ratio of the mean soil solids size, x_n to the maximum aggregate size, X_m at the maximum swelling point; α' , predicted ratio of the mean soil solids size, x_n to the maximum aggregate size, X_{mz} in the oven-dried state; x_m , the maximum sand grain size; $(x_n/X_{mz})_{low}$, lower estimate of the $\alpha' = x_n/X_{mz}$ ratio from soil texture; $(x_n/X_{mz})_{up}$, upper estimate of the $\alpha' = x_n/X_{mz}$ ratio from soil texture.

ence between those and similar values, P_z relating to W=0(Table 1).

Comparing α' (Table 3) with the lower and upper estimates, $(x_n/X_{mz})_{low}$ and $(x_n/X_{mz})_{up}$ (Table 3), one can see that for all the eight soils and three possible $x_{\rm m}$ values the following inequality is fulfilled.

$$(x_n/X_{mz})_{low} < \alpha' < (x_n/X_{mz})_{up} . \tag{26}$$

Thus, all the independently found lower and upper estimates $((x_n/X_{mz})_{low}$ and $(x_n/X_{mz})_{up}$; Table 3) for the soils are in the agreement with the estimates of $\alpha' = x_n/X_{mz}$ (Table 3) that were found from α (Table 3; Eq.(21)), and α , in turn, was found from Eq.(18) with data sets, Y_z , W_h , ρ_s , c, P_z , K, k, and $U_{\rm lpz}$ including the K ratio from Table 1.

These results validate the approach to the K ratio estimation using the additional parameters, X_{mz} and x_n that flow out of the soil aggregate-size distribution in the oven-dried state and soil texture, respectively. Accordingly, the K ratio in the data set of input parameters for prediction of the reference shrinkage curve [1, 2] can be replaced with the $X_{\rm mz}$ and $x_{\rm n}$

It is worth emphasizing that all other input parameters, Y_z , W_h , ρ_s , c, P_z , k, and U_{lpz} as well as two additional parameters, $X_{\rm mz}$ and $x_{\rm n}$ can be found independently of an experimental shrinkage curve. This fact is of principle importance. It means that the model from [1, 2] permits prediction of the soil reference shrinkage curve before its measurement and based only on a number of elementary physical properties of inter- and intraaggregate soil structures.

CONCLUSION

Chertkov [1, 2] introduced the new physical parameter of an aggregated soil - the ratio of an aggregate solid mass to the solid mass of an intraaggregate matrix. Using an experimental shrinkage curve (with a negligible crack volume contribution [1, 2]) this parameter can be easily found as a ratio of two characteristic water content values (without fitting) or as a fitted parameter. This work presents some supplemental consideration and results to [1, 2] to show that the aggregate/intraaggregate mass ratio, as a fundamental property of aggregated soil with any clay content, can be found independently of an experimental shrinkage curve based on a number of measured soil parameters. For that we derive an equation for the aggregate/intraaggregate mass ratio and show that this ratio can be predicted as a root of the equation. This equation is validated using data on eight soils. The measured physical soil parameters for finding the aggregate/intraaggregate mass ratio and predicting the reference shrinkage curve, include: oven-dried specific volume (Y_z) , maximum swelling water content (W_h) , mean solid density (ρ_s) , soil clay content (c), oven-dried structural porosity (P_z) , the lacunar factor that characterizes the rate of the lacunar pore volume change with water content (k), oven-dried lacunar pore volume (U_{lpz}) , the maximum size of aggregates in the oven-dried state (X_{mz}) , and the mean size of the soil solids (clay particles, silt, and sand grains) by their weight fractions (x_n) .

NOTATION

csoil clay content, dimensionless critical clay content, dimensionless

		The Open Hydrology Journal, 2008, Volume 2 13
$F(X,P_{\rm h})$) =	cumulative aggregate-size distribution at a given structural porosity, P_h , dimensionless
$I_{\rm o}(\eta)$	=	function from Eq.(7), dimensionless
K	=	ratio of aggregate solid mass to solid mass of intraaggregate matrix, dimensionless
k	=	lacunar factor, dimensionless
$P_{ m h}$	=	structural porosity of the soil at $W=W_h$, dimensionless
$P_{\rm z}$	=	oven-dried structural porosity, dimension- less
$p_1,,p$	P ₅ =	weight fractions of soil solids from five size ranges (see Table 2), dimensionless
$U_{ m h}$	=	maximum specific volume of the intraaggregate matrix, dm³ kg⁻¹
$U_{ m i}$	=	contribution of interface aggregate layer to the specific volume of soil aggregates, dm ³ kg ⁻¹
$U_{ m lpz}$	=	lacunar pore volume in the oven-dried state, dm³ kg-1
u(W)	=	relative volume of intraaggregate matrix, dimensionless
$u_{\rm h}$	=	relative volume of intraaggregate matrix at $W=W_h$, dimensionless
$u_{\rm s}$	=	relative volume of solids of intraaggregate matrix, dimensionless
u_{z}	=	relative oven-dried volume of intraaggregate matrix, dimensionless
$V_{\rm a}(X)$	=	volume of a separate aggregate of size X , m ³
$V_{\rm i}(X, x)$	=	volume of the interface layer of the mean thickness, x relating to a separate aggregate of size X , m ³
$W_{ m h}$	=	maximum swelling water content, kg kg-1
W	=	soil water content, kg kg ⁻¹
X	=	current aggregate size, m
$X_{\rm m}$	=	maximum size of aggregates, m
X_{\min}	=	minimum size of aggregates, m
$X_{ m mz}$	=	$X_{\rm m}$ in oven-dried state, m
$X_{ m minz}$	=	X_{\min} in oven-dried state, m
x	=	mean thickness of interface layer of aggregates, m
$x_{\rm m}$	=	maximum sand grain size, m
\mathcal{X}_{n}	=	mean size of soil solids, m
$x_{n \text{ low}}$	=	lower estimate of the mean solids size, x_n , m
$x_{\text{n up}}$	=	upper estimate of the mean solids size, x_n , m

upper estimate of the x_n/X_{mz} ratio, dimen-

lower estimate of the x_n/X_{mz} ratio, dimen-

oven-dried specific volume of soil, dm³ kg⁻¹

 $(x_n/X_{mz})_{up}=$

 $(x_n/X_{mz})_{low} =$

 $Y_{\rm z}$

sionless

 α = ratio from Eq.(19), dimensionless

 α' = ratio from Eq.(21), dimensionless

η = parameter from Eq.(6), dimensionless

 ρ_b = dry bulk density, g cm⁻³

 ρ_s = mean density of solids, g cm⁻³

REFERENCES

- Chertkov VY. The reference shrinkage curve at higher than critical soil clay content. Soil Sci Soc Am J 2007; 71(3): 641-655.
- [2] Chertkov VY. The soil reference shrinkage curve. Open Hydrol J 2007: 1: 1-18.
- [3] Braudeau E, Frangi J-P, Mohtar RH. Characterizing nonrigid aggregated soil-water medium using its shrinkage curve. Soil Sci Soc Am J 2004; 68: 359-370.
- [4] Peng X, Horn R. Modeling soil shrinkage curve across a wide range of soil types. Soil Sci Soc Am J 2005; 69: 584-592.
- [5] Cornelis WM, Corluy J, Medina H, Hartmann R, Van Meirvenne M, Ruiz ME. A simplified parametric model to describe the magni-

- tude and geometry of soil shrinkage. Eur J Soil Sci 2006; 57: 258-268.
- [6] Cornelis WM, Corluy J, Medina H, et al. Measuring and modelling the soil shrinkage characteristic curve. Geoderma 2006; 137: 179-191.
- [7] Chertkov VY. Modeling the pore structure and shrinkage curve of soil clay matrix. Geoderma 2000; 95: 215-246.
- [8] Chertkov VY. Modelling the shrinkage curve of soil clay pastes. Geoderma 2003; 112: 71-95.
- [9] Reeve MJ, Hall DGM. Shrinkage in clayey subsoils of contrasting structure. J Soil Sci 1978; 29: 315-323.
- [10] Baer JU, Anderson SN. Landscape effects on desiccation cracking in an Aqualf. Soil Sci Soc Am J 1997; 61: 1497-1502.
- [11] Braudeau E, Sene M, Mohtar RH. Hydrostructural characteristics of two African tropical soils. Eur J Soil Sci 2005; 56: 375-388.
- [12] Chertkov VY. The reference shrinkage curve of clay soil. Theor Appl Fract Mech 2007; 48(1): 50-67.
- [13] Chertkov VY. Intersecting-surfaces approach to soil structure. Int Agrophysics 2005; 19: 109-118.
- [14] Hillel D. Environmental Soil Physics. San Diego Academic Press

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