

Atmospheric tide disturbances as Earthquake precursory phenomena

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Abstract. The tidal changes of the barometric pressure in the area of Thessaloniki were studied by analysing a sample of 21 years of hourly measurements. The resulted tidal parameters (amplitude and phase difference) were considered as “mean values” of the corresponding parameters for this long time period. Using these parameters, barometric changes were computed and subtracted from the observations. Assuming that the residuals might include local (in terms of time) information, the residual time series was split in 21 blocks and a new analysis was performed for each block separately. The 21 amplitude values computed for each tidal wave were considered as the amplitude variation with respect to the corresponding mean value. An observable correlation of the amplitude exalting of the most of the tidal waves was found to Earthquakes of magnitude >4 , occurred close to Thessaloniki in the test period.

1 Introduction

It is well-known that the changes in the barometric pressure influence the measurement of the gravity field and a great deal of research is directed to the determination of this influence (see, Merriam, 1992).

Theoretical calculation based on mathematical models lead to a value of $-0.43 \times 10^{-8} \text{ ms}^{-2}/\text{hPa}$. In fact, lithospheric and water surface deformations due to barometric pressure changes reduce this influence to -0.2 to $-0.3 \times 10^{-8} \text{ ms}^{-2}/\text{hPa}$ (Deike, 1975).

Although the effect of the atmospheric tides is much weaker than that of the oceanic tides, it has been found that atmospheric tides are mainly responsible for the noise to the determination of the local potential (Warburton and Goodkind, 1977). So the knowledge of the atmospheric tides at a site became of particular interest.

Atmospheric tides have been a subject of extensive theoretical and observational research during the last two centuries, so that today we possess an adequate theoretical explanation of the observed phenomena. An extensive presentation of the state-of-the-art up to the recent years is given by Chapman and Lindzen (1970) while an extensive analysis of the diurnal and semidiurnal thermogravitational solar tide of the Earth atmosphere in the area of Greece was performed by Barzokas (1989).

In addition to the diurnal and semidiurnal lunisolar tidal constituents, S1, S2, M1 and M2, whose effect on the Earth atmosphere is well documented, there are also the weaker lunisolar tidal constituent of lower frequencies, whose effect on the Earth atmosphere has not been subject to detailed research until today.

However, the marginal presence of the frequencies 0.03629 and 0.073036 cycles/day in the power spectrum of the daily measurements series of barometric pressure was detected during an investigation upon the influence of the barometric pressure and the Earth tides on the shallow underground water level variations in the area of Volvi, Northern Greece. These frequencies coincide with those of the Mm (period of 27.55 days) and Mf (period of 13.69 days) constituents of the Earth tides indicating a possible influence of the Earth tides on barometric pressure (Contadakis, 1994).

In an earlier paper (Arabelos et al., 1997) an analysis of a sample of 11.75 years of hourly measurements (103 032 values) showed a clear tidal response at Sa, Ssa, S1 and S2, which corresponds to the solar annual, semiannual, diurnal and semidiurnal constituents. In the same analysis the response at P1, K1 and K2 was detectable, while the diurnal M1, semidiurnal M2 and terdiurnal M3 lunar constituents failed to be identified since the value of the corresponding signal-to-noise ratio was very low.

The present work was inspired from a recent paper presented at the last EGS General Assembly (Biagi, et al., 2003). In this paper low frequency radio signals were analysed, in order to reveal exalting in atmospheric tides. The exalting found was correlated with a strong seismic crisis in Slovenia

during the following months, since the preparatory phase of Slovenia Earthquakes could have produced such a local exalting of the atmospheric tides. According to the authors, the possible explanation is the following: The preparatory phase of the seismic crisis produced a diffusion of fluids in a broad area around the Earthquake epicentres. Such a diffusion of fluids was observed many times before Earthquakes. As a consequence, the ground density in this zone decreased, resulting in a local decrease of gravity. Due to the gravity decrease, the lower atmosphere overlying the perturbed zone expands and its density decreases. In consequence of this density decrease the local atmospheric tides can be amplified.

In the present paper it was attempted to reveal exalting in atmospheric tides by analyzing barometric pressure measurements. A 21 years data set of hourly values was used for this analysis. In order to exclude that the atmospheric tide disturbances are due to changes of the atmospheric temperature close to the Earth's surface, the correlation between atmospheric tides and atmospheric temperature was estimated using a similar data set of hourly atmospheric temperature values. A linear regression model was used for this purpose. Then the correlated part was subtracted from the original observations.

The analysis of such a long period data set provided tidal parameters that could be considered as "mean values", describing the general behaviour of atmospheric tides in the area. In fact there are not considerable deviations between the estimated amplitudes in the present analysis and the corresponding ones, found from the analysis of the 11.75 years of hourly measurements by Arabelos et al. (1997).

Using these parameters, hourly atmospheric tides were computed for the same period and subtracted from the observations. Assuming that the residuals might include local (in terms of time) information, the residual data set was split in 21 blocks and a new analysis was carried out for each block separately. The 21 amplitude values computed for each tidal wave were considered as the amplitude variation with respect to the corresponding "mean value" mentioned previously.

The advantage in this procedure is that we are able to have a detailed analysis for all waves with periods from 365.2568 to 0.2588 days. On the other hand, in order to get reliable information concerning the annual waves such as S_a , it is necessary to analyse time series covering a period at least of one year. In this way, the resolution of this method is restricted. In fact, the estimation of the amplitude for the annual solar waves in our analysis of the yearly blocks resulted in relatively large estimation errors with respect to the corresponding error for the parameters of lower-period waves.

2 Observational material and model of normal potential

In the present analysis the data sets used were the following:

- Model of tidal potential

From the models of Normal Tidal Potential we chose that of Tamura (1987), which comprises 1200 waves.

- Barometric pressure data

The atmospheric pressure has been recorded with the help of a membrane box barograph and corrected with the help of a mercury pipe barometer installed at the same place with the barograph. The accuracy of each measurement of pressure is ± 0.1 hPa.

The data set includes 184 080 hourly values of the barometric pressure covering the period from 1 January 1981 to 31 December 2001 (totally 7670 days). These values have been derived from analog recordings of the meteorologic station of the Department of Meteorology and Climatology of the Aristotle University of Thessaloniki. The data set was used for the determination of the atmospheric tidal parameters. The statistical characteristics of these values are given in Sect. 3.

- Atmospheric temperature data

The atmospheric temperature was recorded with the help of a bimetallic thermograph and corrected with the help of a mercury thermometer. The accuracy of each measurement of temperature is $\pm 0.1^\circ\text{C}$. The hourly values of the temperature for the same period as in the section on Barometric pressure data have been used to estimate the correlation between atmospheric temperature close to the Earth's surface and barometric pressure at the tidal frequencies. The statistical characteristics of these values are given in Sect. 3.

- Earthquake catalogue

The catalogue of the Geodynamic Institute/National Observatory of Athens was used, comprising 43 128 Earthquakes from 1964 up to now, which occurred in the area bounded by $33.47^\circ < \varphi < 42.20^\circ$, $16.26^\circ < \lambda < 30.36^\circ$. The magnitude of these events varies from 1.4 to 7.

3 Analysis of the barometric measurements

In order to detect tidal signal in the changes of the barometric pressure the following analysis of the barometric measurements has been performed. The barometric measurements were compared with the normal values of the tidal potential and the residuals are used in order to estimate the tidal parameters of the barometric tides through a least squares procedure. These parameters can be used to predict the barometric tides in the same station at any time. The method of the analysis was similar to that used for the analysis of the gravimetric tides (e.g. Arabelos, 2002) or the ocean tides (e.g. Spatalas, 1995). There is a difference in relation to the gravimetric tides: The corresponding measurements are usually affected by the instrumental drift. Consequently, an appropriate filtering procedure is necessary in order to remove the drift effect from the corresponding time series. This filtering

Table 1. Statistics of the barometric measurements and air temperature measurements.

| Year | Time span (hours) | Barometric pressure | | Air temperature | |
|-------|----------------------|---------------------|--------------------------------|--------------------|-------------------------------|
| | | Mean value (hPa) | Standard deviation (hPa) | Mean value (°C) | Standard deviation (°C) |
| 1981 | 8760 | 1018.409 | 6.545 | 16.09 | 7.93 |
| 1982 | 8760 | 1020.542 | 6.498 | 15.65 | 7.28 |
| 1983 | 8760 | 1014.698 | 6.990 | 15.43 | 8.11 |
| 1984 | 8784 | 1013.344 | 6.134 | 15.57 | 7.78 |
| 1985 | 8760 | 1013.323 | 6.608 | 16.12 | 8.80 |
| 1986 | 8760 | 1014.788 | 5.958 | 15.39 | 8.26 |
| 1987 | 8760 | 1017.949 | 6.190 | 15.28 | 8.87 |
| 1988 | 8784 | 1018.232 | 6.089 | 15.55 | 8.88 |
| 1989 | 8760 | 1020.095 | 6.946 | 15.55 | 7.88 |
| 1990 | 8760 | 1019.528 | 6.636 | 16.20 | 8.06 |
| 1991 | 8760 | 1019.492 | 5.699 | 15.10 | 8.55 |
| 1992 | 8784 | 1020.230 | 6.420 | 15.54 | 8.28 |
| 1993 | 8760 | 1014.033 | 7.104 | 15.61 | 8.53 |
| 1994 | 8760 | 1012.958 | 5.574 | 16.69 | 8.10 |
| 1995 | 8760 | 1016.131 | 6.228 | 15.90 | 8.24 |
| 1996 | 8784 | 1015.834 | 5.838 | 15.21 | 8.19 |
| 1997 | 8760 | 1017.153 | 6.391 | 15.29 | 8.30 |
| 1998 | 8760 | 1016.904 | 6.047 | 16.04 | 8.62 |
| 1999 | 8760 | 1016.270 | 6.044 | 16.46 | 8.38 |
| 2000 | 8784 | 1017.760 | 6.081 | 16.46 | 8.62 |
| 2001 | 8760 | 1016.157 | 6.281 | 16.68 | 8.82 |
| Total | 18 4080 | 1016.850 | 6.733 | 15.87 | 8.37 |

corresponds to an amplification of the tidal frequencies and an attenuation of the drift frequencies (Wenzel, 1976). During this filtering procedure it is dangerous to interpret a part of the tidal signal as drift and vice versa. In the case of barometric measurements it is not necessary to take care for the estimation of the drift and the corresponding correction since the barograph is considered to be a drift-free instrument, at least for the time periods used in this analysis.

The time series covering 21 years of barometric pressure registrations enabled us to include also long period waves in the analysis. Specifically, parameters for 22 waves were estimated, covering periods from 365.2568 to 0.345 days.

A significant step in the procedure of the analysis is the test of the accuracy of the observations. For a short time period the observed signal could be approximated by the function (Wenzel, 1976)

$$\begin{aligned}
 y(t) = & a_1 + a_2 t + a_3 t^2 + a_4 \cos 15^\circ t + a_5 \sin 15^\circ t \\
 & + a_6 \cos 30^\circ t + a_7 \sin 30^\circ t + a_8 \cos 45^\circ t \\
 & + a_9 \sin 45^\circ t.
 \end{aligned} \quad (1)$$

The values of the coefficients a_1, \dots, a_9 are defined through an adjustment procedure substituting in Eq. (1) the values $t = -11^h, \dots, 11^h, t \neq 0$. The deviation between the observation corresponding to $t = 0$ and the approximation

$$D = L_{(0)} - \tilde{y}_{(0)} = L_{(0)} - (\bar{a}_1 + \bar{a}_4 + \bar{a}_6 + \bar{a}_8) \quad (2)$$

could be used as a criterion for the existence of errors in the observations. Equation (2) describes the numerical filtering of the measurements

$$D = \mathbf{f}_l^T \mathbf{L}, \quad (3)$$

where \mathbf{f}_l^T is the vector of the coefficients of the numerical filter and \mathbf{L} the vector of the observations

$$L_{(t)}, t = -11, \dots, 0, \dots, 11.$$

Applying the above described procedure and adopting the value $D = 5$ hPa, a number of outliers was detected arising from digitization errors. After the correction of these values the statistics of the data is shown in Table 1.

From Table 1 it is evident that the annual standard deviation of the barometric pressure in Thessaloniki was almost stable during the period of the twenty one years with the exception of 1986, 1991, 1994 and 1996.

The FORTRAN program ETERNA (Wenzel, 1996) has been used for the analysis. This program allows the adjustment of tidal parameters from either unfiltered or highpass filtered observations and the computation of the FOURIER-amplitude spectrum of the residuals. When using unfiltered observations, tidal parameters from long- to short-periodic waves may be adjusted. Also, the program computes the error estimation of the adjusted tidal parameters: (a) by a

Table 2. Adjusted tidal parameters of the barometric field in the area of Thessaloniki, taking into account the correlation between pressure and air temperature. Error estimation by FOURIER-spectrum of residuals. In parentheses is the error estimation by least squares method.

| Waves | Period (days) | Signal-to-noise ratio | Amplitude (hPa) | Phase Difference (degree) |
|-------|---------------|-----------------------|-----------------------------|------------------------------|
| Sa | 365.2568 | 19.2(84.6) | $2.538 \pm 0.10403(.01542)$ | $105.622 \pm 2.989(2.028)$ |
| Ssa | 182.6217 | 2.4(15.7) | $0.312 \pm 0.01652(.00281)$ | $101.832 \pm 24.278(.141)$ |
| MM | 27.5545 | 3.2(23.0) | $0.418 \pm 0.01456(.00158)$ | $-18.136 \pm 18.129(.115)$ |
| MF | 13.6608 | 1.9(13.7) | $0.248 \pm 0.00769(.00083)$ | $-132.482 \pm 30.651(.060)$ |
| MTM | 9.1329 | 1.1 (8.4) | $0.145 \pm 0.04015(.00416)$ | $-148.640 \pm 52.203(.302)$ |
| Q1 | 1.1195 | 0.3 (.2) | $0.005 \pm 0.00022(.00025)$ | $-120.018 \pm 165.947(.018)$ |
| O1 | 1.0758 | 1.5 (1.0) | $0.020 \pm 0.00004(.00005)$ | $-31.309 \pm 37.716(.004)$ |
| M1 | 1.0347 | 0.4 (.3) | $0.005 \pm 0.00054(.00058)$ | $122.900 \pm 145.515(.042)$ |
| P1 | 1.0027 | 17.4(11.5) | $0.227 \pm 0.00009(.00015)$ | $-176.446 \pm 3.290(.008)$ |
| S1 | 1.0000 | 113.9(43.5) | $1.486 \pm 0.00386(.01181)$ | $-46.456 \pm 0.503(.576)$ |
| K1 | .9973 | 19.0(12.6) | $0.248 \pm 0.00003(.00006)$ | $145.199 \pm 3.012(.003)$ |
| PSI1 | .9946 | 4.2 (2.8) | $0.055 \pm 0.00386(.00747)$ | $158.276 \pm 13.511(.330)$ |
| PHI1 | .9919 | 3.2 (2.1) | $0.041 \pm 0.00212(.00344)$ | $41.209 \pm 18.047(.185)$ |
| J1 | .9624 | 0.9 (.6) | $0.011 \pm 0.00054(.00062)$ | $102.921 \pm 66.658(.045)$ |
| OO1 | .9294 | 0.8 (.7) | $0.011 \pm 0.00099(.00094)$ | $20.827 \pm 69.559(.068)$ |
| 2N2 | .5363 | 0.3 (.1) | $0.002 \pm 0.00050(.00089)$ | $-35.769 \pm 227.108(.064)$ |
| N2 | .5274 | 0.5 (.2) | $0.004 \pm 0.00008(.00018)$ | $141.364 \pm 107.726(.013)$ |
| M2 | .5175 | 2.5 (.8) | $0.017 \pm 0.00002(.00004)$ | $169.438 \pm 23.007(.003)$ |
| L2 | .5080 | 0.5 (.2) | $0.004 \pm 0.00055(.00109)$ | $42.840 \pm 108.577(.079)$ |
| S2 | .5000 | 69.5(23.5) | $0.464 \pm 0.00003(.00008)$ | $-154.667 \pm 0.824(.006)$ |
| K2 | .4986 | 9.0 (3.2) | $0.060 \pm 0.00012(.00027)$ | $167.962 \pm 6.358(.020)$ |
| M3 | .3450 | 0.4 (.1) | $0.002 \pm 0.00084(.00228)$ | $55.980 \pm 137.304(.166)$ |

least squares adjustment procedure and (b) by FOURIER-amplitude spectrum of the residuals of the adjustment, using average spectral amplitudes over individual tidal bands. The last method is known to give more realistic error estimations, especially when adjusting simultaneously long- and short-periodic tidal parameters (Wenzel, 1976) as in our case.

As it was mentioned in Sect. 1 the correlation between barometric pressure and air temperature was estimated, using a linear regression model. The analysis showed a relatively weak correlation with a regression coefficient equal to -0.36864 ± 0.00420 hPa/ $^{\circ}$ C being obtained. However, the last result is still significant. This was shown through the following numerical experiment. The analysis of the entire time series was carried out in two ways: (i) ignoring the correlation and (ii) taking into account this correlation by removing from the observations the correlated part. The results showed that the waves mainly affected from the temperature are the solar Sa (15% amplitude decrease), Ssa (107% amplitude increase), S1 (142% amplitude increase) and S2 (21% amplitude increase). Significant changes also appeared in the case of the diurnal waves P1 and K1.

In Table 2 the results of the analysis of the barometric pressure measurements for the 21 years period is shown. The correlation between barometric pressure and air temperature was taken into account and the error estimation was computed from the FOURIER-spectrum of the residuals. The error estimation by the least squares method is also shown in parentheses.

The standard deviation of the weight unit in this adjustment was 0.973 hPa. There is a very good agreement with corresponding results computed from the analysis of an 11.75 years data set (Arabelos et al., 1997). This is a confirmation that the estimated amplitudes are characteristic of the atmospheric tides in the test area.

Using the values of the tidal parameters of Table 2, the tidal changes of the barometric field during the entire period were computed and subtracted from the observations. Subsequently, the resulted residual data set was split in 21 blocks and a new analysis was carried out for each block separately. The computed 21 “residual amplitude” values for each tidal wave were considered as the amplitude variation of each wave, with respect to the corresponding “mean value” of Table 2.

In order to examine if these variations include some information related to the preparatory phase of seismic events, we have split the tidal waves in three groups according to their period and plotted the residual amplitudes together with the seismic events retrieved from the Earthquake catalogue mentioned in Sect. 2. However, only the waves presenting a signal to noise ratio >8 (according to least squares estimation of Table 2) were included in these plots. In Figs. 1, 2 and 3 the annual variation of the residual amplitudes for the various wave groups is shown, together with the Earthquakes of $M \geq 4$ up to a distance of 200 km from Thessaloniki. The real-time of occurrence of the Earthquakes is expressed as Julian Date (J.D.) since 1980, and their focal depths are shown

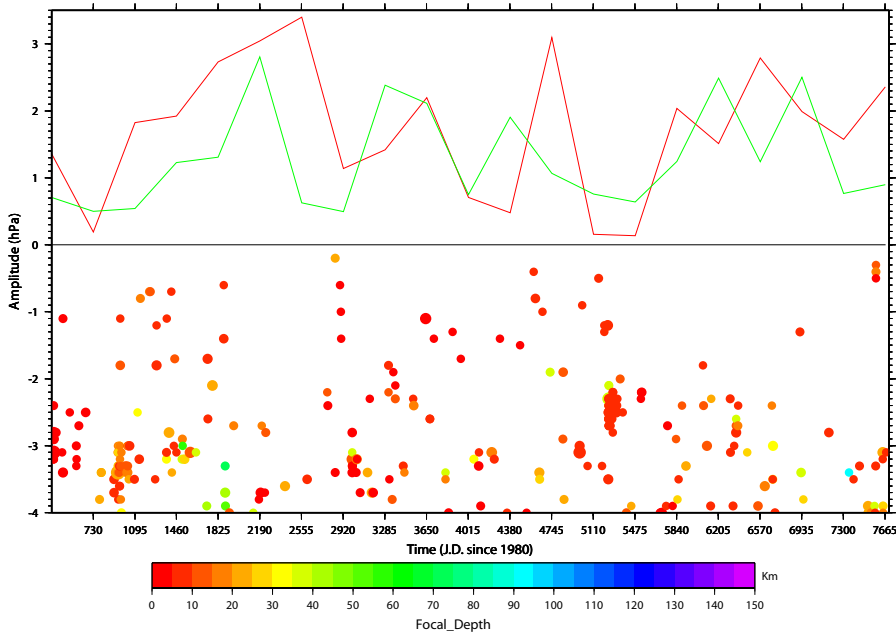


Fig. 1. Upper part: Annual amplitude variation of Sa (red) and Ssa (green). Lower part: Earthquakes that occurred during the test period, up to a distance of 200 km from Thessaloniki.

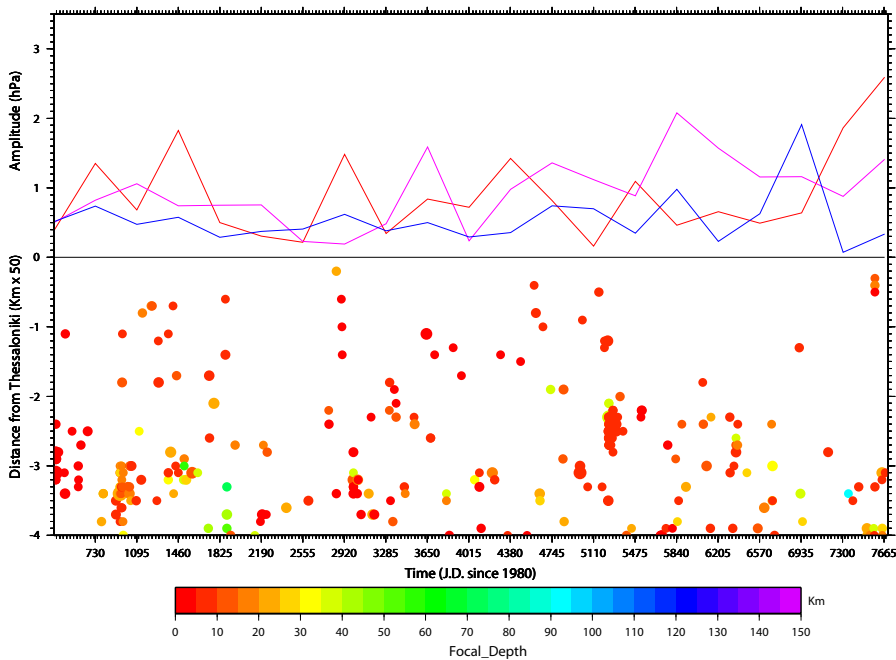


Fig. 2. Upper part: Annual amplitude variation of Mm (red), Mf (magenta) and Mtm (blue). Lower part: Earthquakes that occurred during the test period, up to a distance of 200 km from Thessaloniki.

through the coloured scale. The magnitude of the Earthquake is expressed by the radius of the corresponding circle.

However, there is a problem in comparing a single amplitude value computed from the analysis of a year-long time series with the flowing, real-time occurrence of the seismic events. This problem is more serious in the case of the annual waves, because here it is necessary to analyze observations of at least one year in duration. A possible solution might be to present monthly release of energy, instead of just Earthquake count. But in the time period considered, there are only few Earthquakes with $M \geq 5$. In Figs. 1, 2 and 3 the computed annual amplitudes were plotted at the beginning of each year.

In Figs. 1, 2 and 3 it is shown that the amplitudes of Sa and Ssa waves present an increase correlated with most of the Earthquakes that occurred near Thessaloniki, with some exceptions. It is characteristic the exalting of both waves in the case of the seismic sequence close to J.D. 2920. However, there are cases such as the seismic sequence at the end of J.D. 7665, where the amplitude of Ssa was increased, while the amplitude of Sa was decreased. But it should be mentioned again, that the accuracy estimation of the Sa amplitudes is relatively lower than the accuracy estimation for the other waves, as it was explained in Sect. 1.

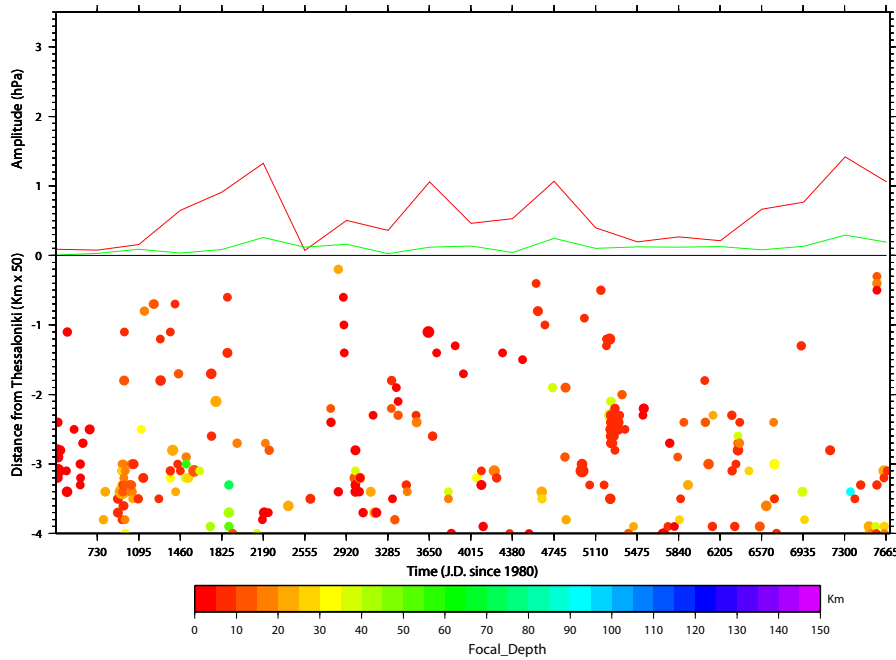


Fig. 3. Upper part: Annual amplitude variation of S1 (red) and S2 (green). Lower part: Earthquakes that occurred during the test period, up to a distance of 200 km from Thessaloniki.

The situation in the case of the moon-monthly, semi-monthly and ter-monthly waves is rather confusing (see Fig. 2). In both cases of Figs. 1 and 2 exalting is shown in the period from about J.D. 5290 to 6570, without the occurrence of any Earthquakes, at least close to Thessaloniki.

More clear is the behaviour of the S1 amplitude (Fig. 3) presenting a systematic increase before the most of the Earthquakes, with a very few exceptions. On the other hand, the amplitude of the semi-diurnal constituents does not show any correlation with the seismic events, even for those occurred very close to Thessaloniki.

4 Conclusions

The tidal analysis of the twenty one years of hourly barometric pressure measurements, showed a clear tidal response at Sa, Ssa, S1 and S2, and a detectable response at P1, K1 and K2 constituents. In the same analysis, the response of the lunar Mm, Mf and Mtm is marginal, while the lunar diurnal M1, semidiurnal M2 and terdiurnal M3 failed to be identified. The results of this analysis agree very well with the results of an earlier analysis of a sample of 11.75 years. This is a confirmation that the estimated amplitude and phase difference parameters of the atmospheric tides are characteristic for the area of Thessaloniki.

An attempt to estimate the correlation between barometric pressure and air temperature using a linear regression model resulted in the estimation of a correlation coefficient equal to -0.36864 ± 0.0042 hPa. The correlated part was subtracted from the original observations, leading in the estimation of tidal parameters independent from the air temperature variations.

The analysis of the residual signal separated in one-year blocks, showed considerable amplitude variations that might include information related to other phenomena, such as Earthquakes. A preliminary attempt in this direction showed that in many cases, an amplitude exalting of some waves may precede an Earthquake close to the test area. Such a behaviour is more evident in the case of S1, but is also observable in the case of Sa and Ssa. However, this does not prove the existence of a relationship between atmospheric tide disturbances and Earthquakes. It merely demonstrates that such a relationship cannot be ruled out, at least without additional meticulous research.

A more detailed analysis aiming to the estimation of parameters from blocks covering shorter time intervals could be very helpful in concluding if the exalting of the tidal waves could be considered as Earthquake precursors.

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